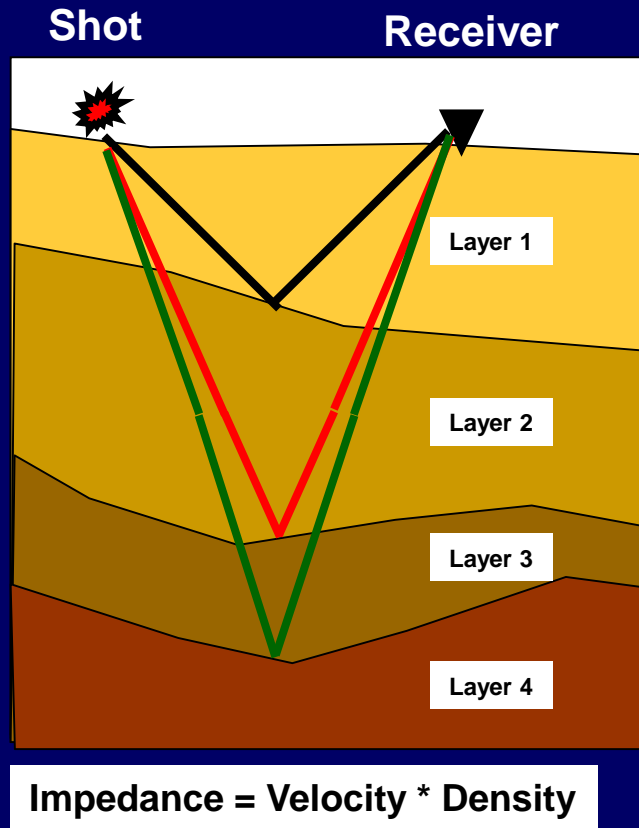
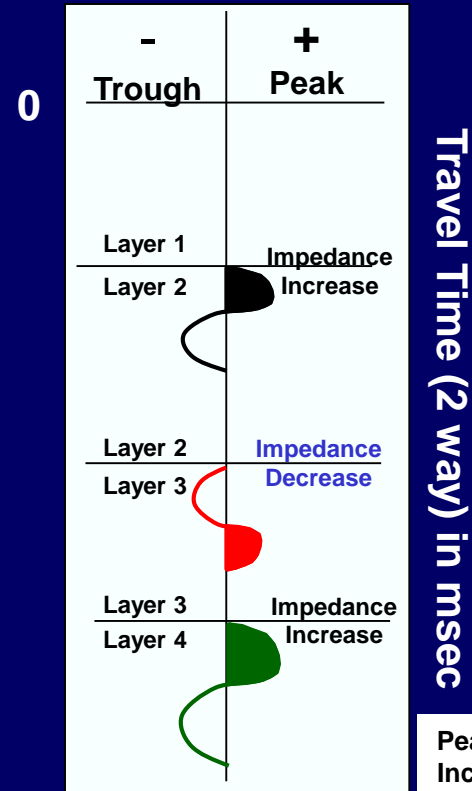


Seismic Reflections

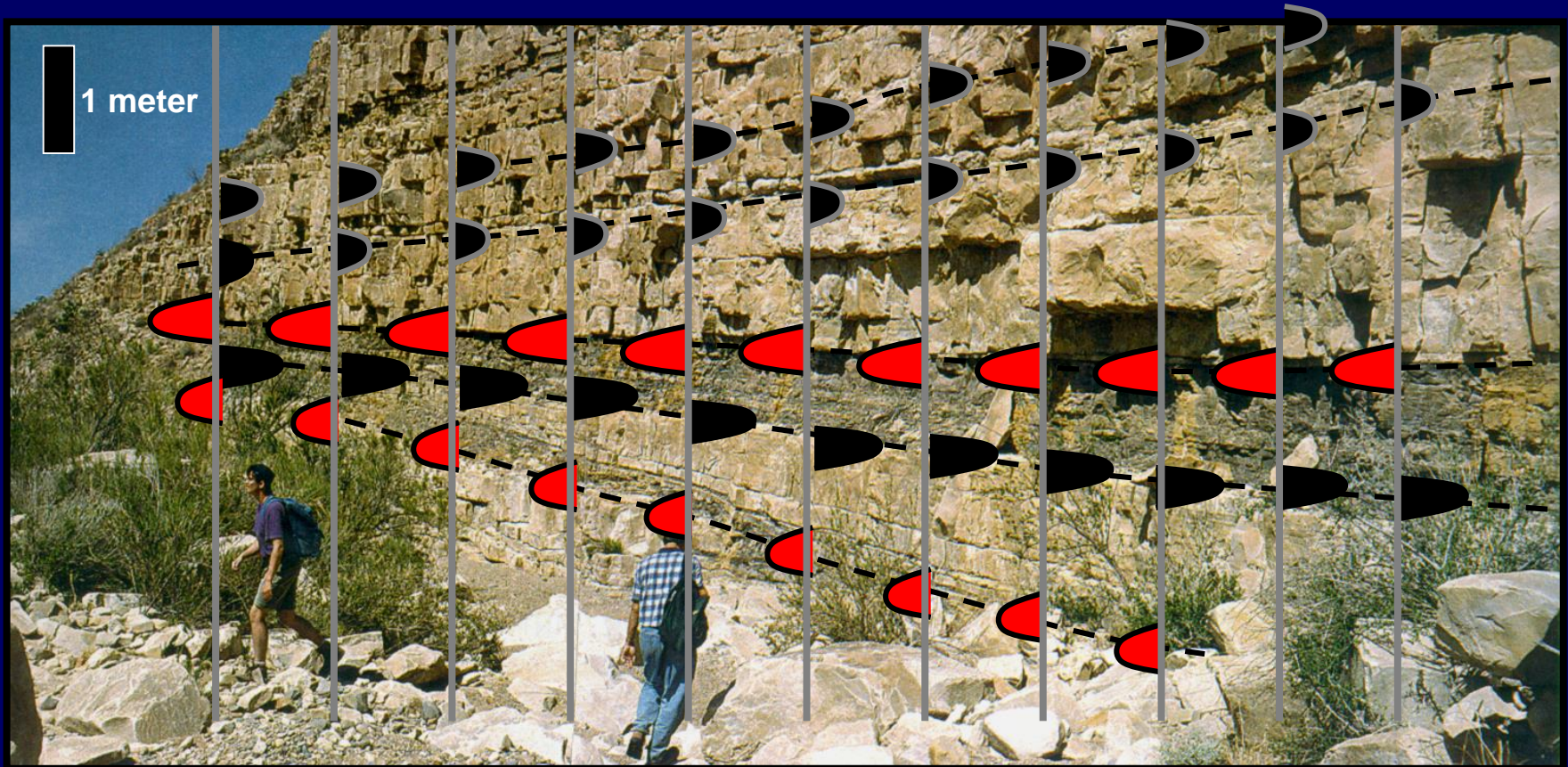


Seismic Record

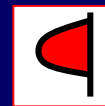


The Ideal Seismic Response

Able to resolve boundaries of beds a few meters thick

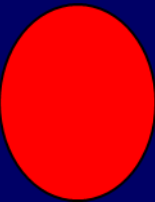


Increase in Impedance



Decrease in Impedance

Scale for Seismic Data



Lamina

Lamina Sets

Beds

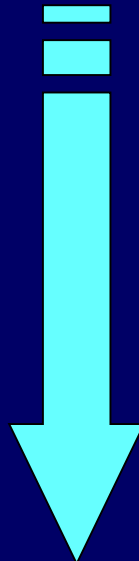
Bed Sets

Parasequences

Parasequence Sets

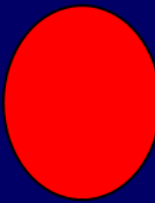
Sequences

Sequence Sets

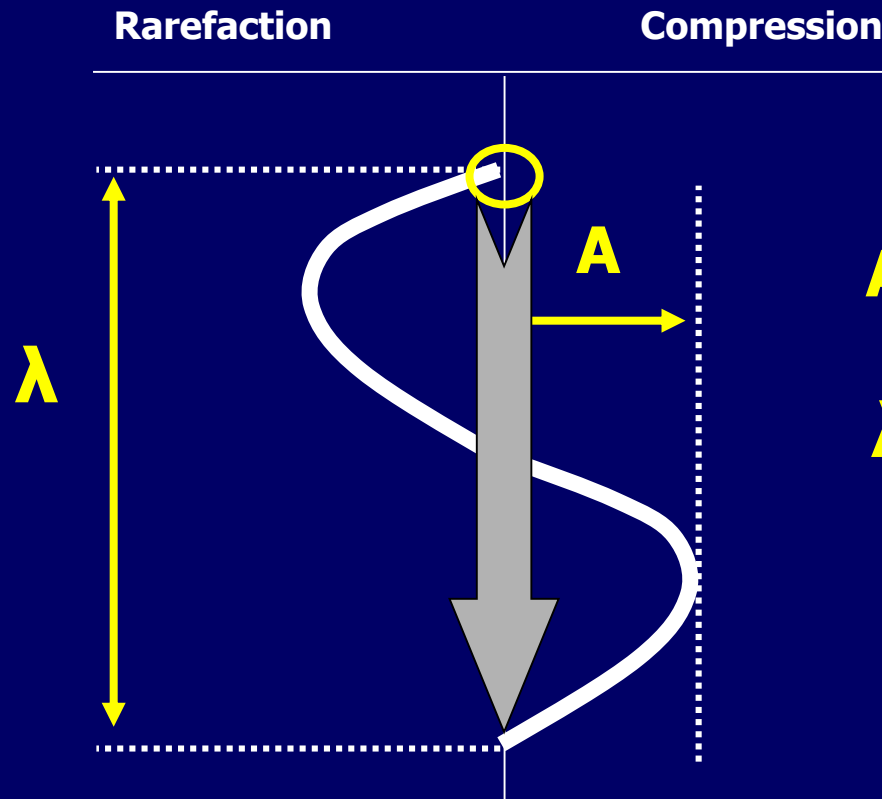


- Although seismic data can not image small-scale stratal units, it can image mid- to large-scale units
- The big advantage of seismic data is areal coverage

Seismic - Units 10s of Meters Thick



Wave Equation Lingo



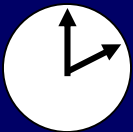
A = Amplitude

λ = Wavelength
length, ft or m

P = Period
time

**D_p = Pulse
Duration**
time

**Period = Time for the waveform
to travel 1 wavelength**



Basic Equations

$$1. P = 1 / f$$

$$2. \lambda = V * P = V / f$$

$$3. d = V * T / 2$$

where

P = Period

V = Velocity

f = Frequency

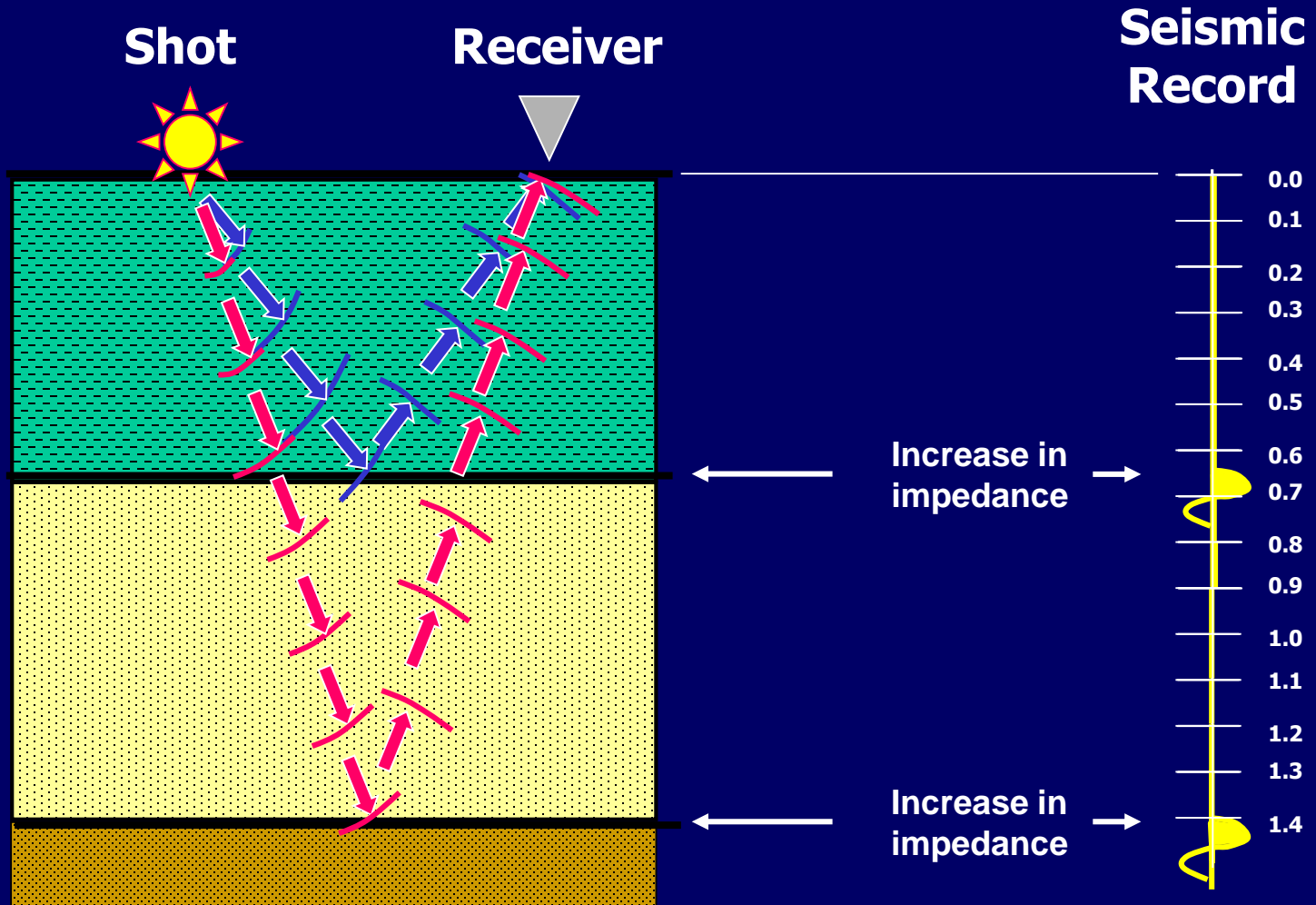
d = distance (depth)

λ = Wavelength

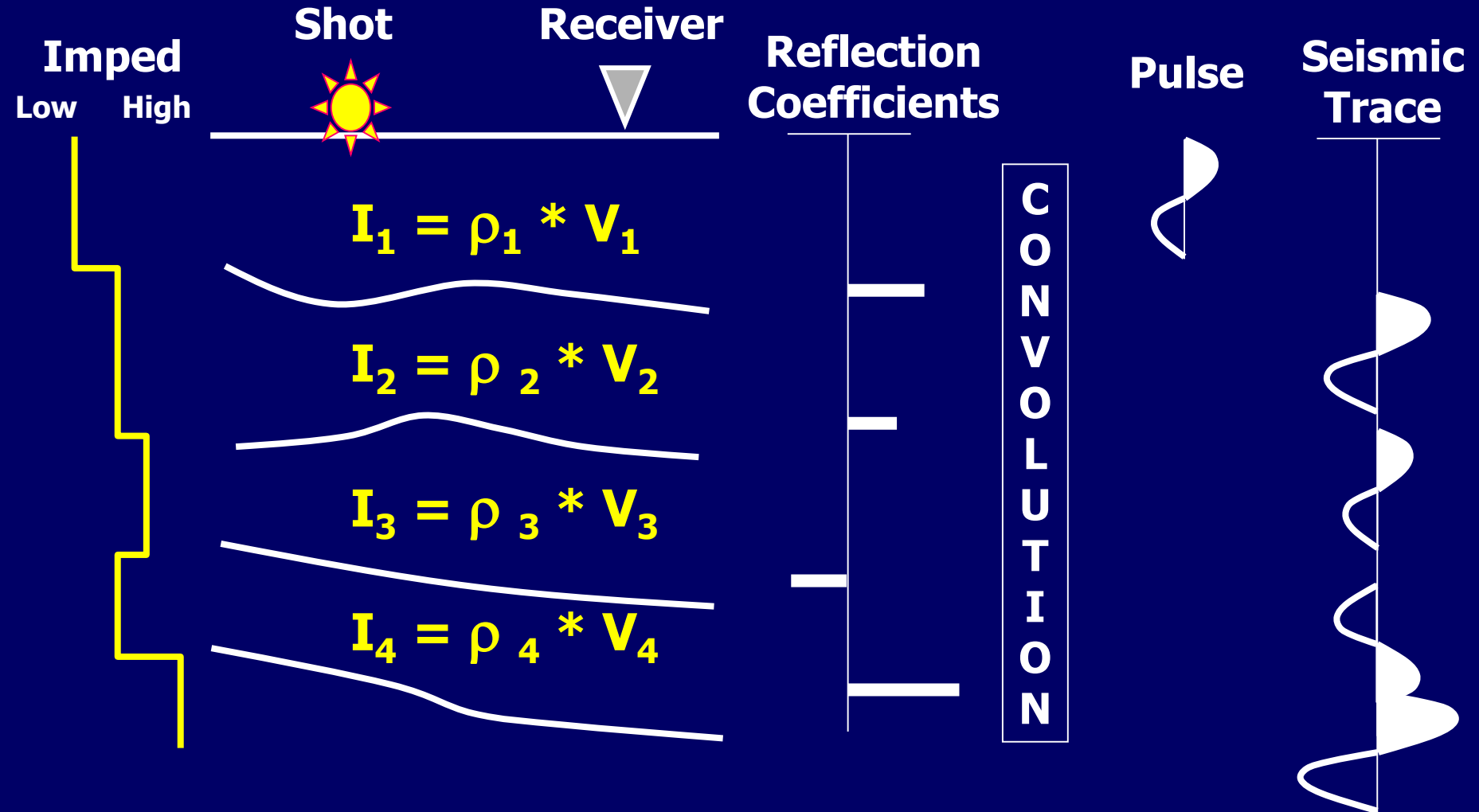
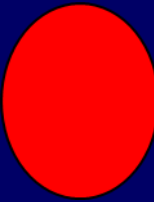
T = time

Back to Basics

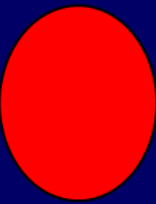
Seismic energy travels down and is reflected off acoustic boundaries



Acoustic Structure of the Earth



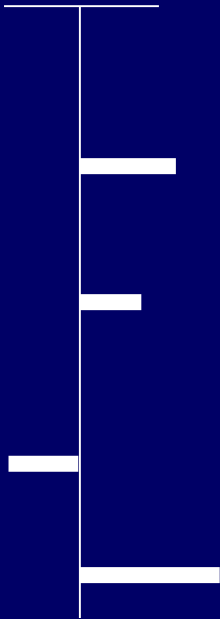
That 'Pesky' Pulse



If the frequency content (Bandwidth) is very large, then the pulse approaches a spike and we can resolve fine-scale stratigraphy

Typically the frequency content is limited to about 10 to 50 Hz ($BW = 40$), which limits our resolution

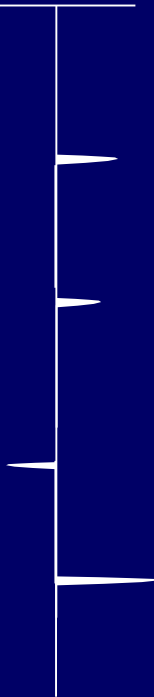
Reflection Coefficients



Ideal Pulse



Seismic Trace



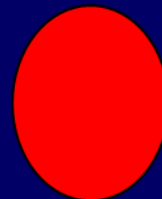
Typical Pulse



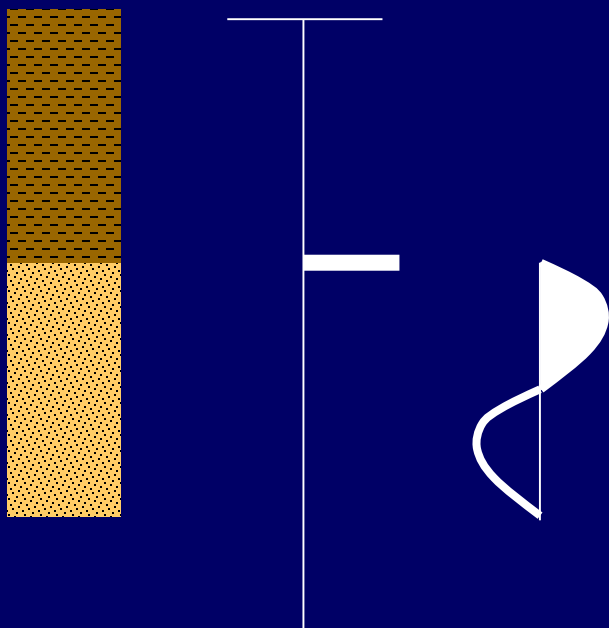
Seismic Trace



Types of Pulses



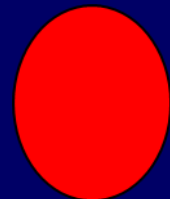
Reflection Coefficients



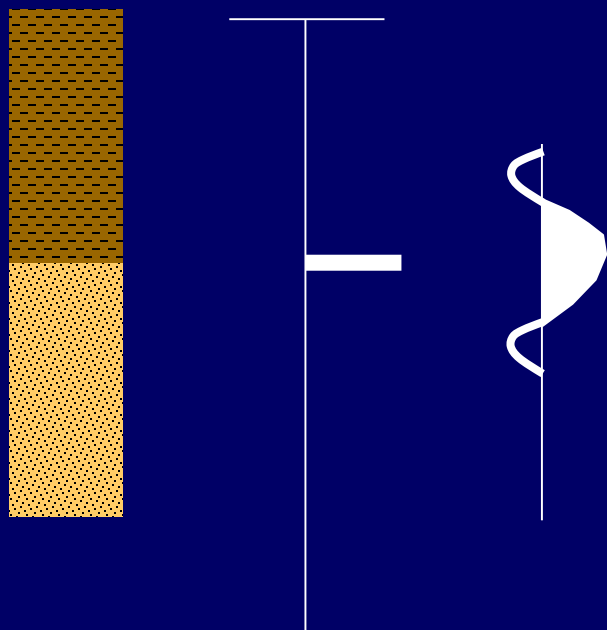
Minimum Phase

- Causal (real – no motion before wave arrives)
- Front loaded
- Peak arrival time is frequency dependent
- RC is at the first displacement; maximum displacement (peak or trough) is delayed by $\frac{1}{4} \lambda$

Types of Pulses



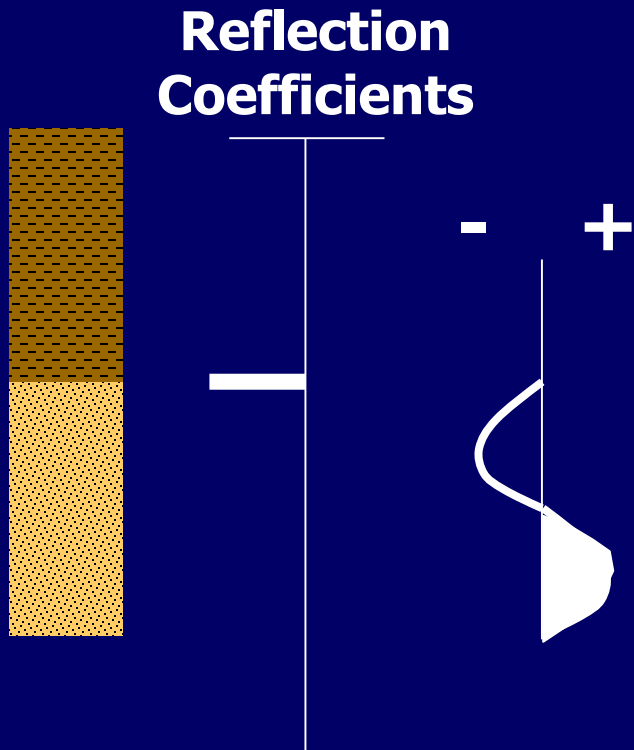
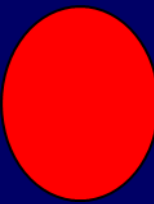
Reflection Coefficients



Zero Phase

- Not Causal (not real, since there is motion before the wave arrives)
- Symmetric about RC
- Peak arrival time is not frequency dependent
- Maximum peak-to-side lobe ratio
- RC is at the maximum displacement (peak or trough)

Polarity – Minimum Phase



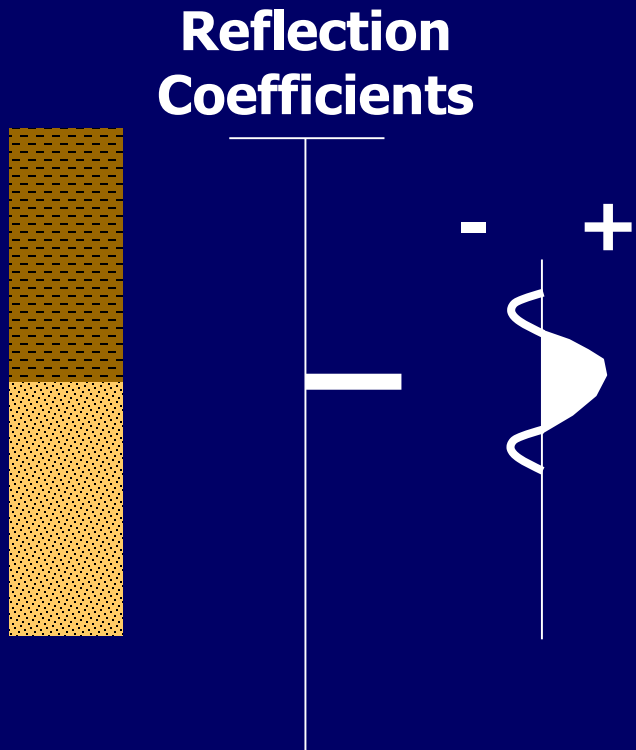
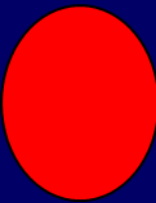
SEG Normal Convention

A compression is:

- Negative # on the tape
- Displayed as a Trough

SEG = Society of Exploration Geophysics

Polarity – Zero Phase



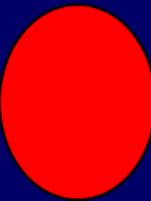
SEG Normal Convention

A compression is:

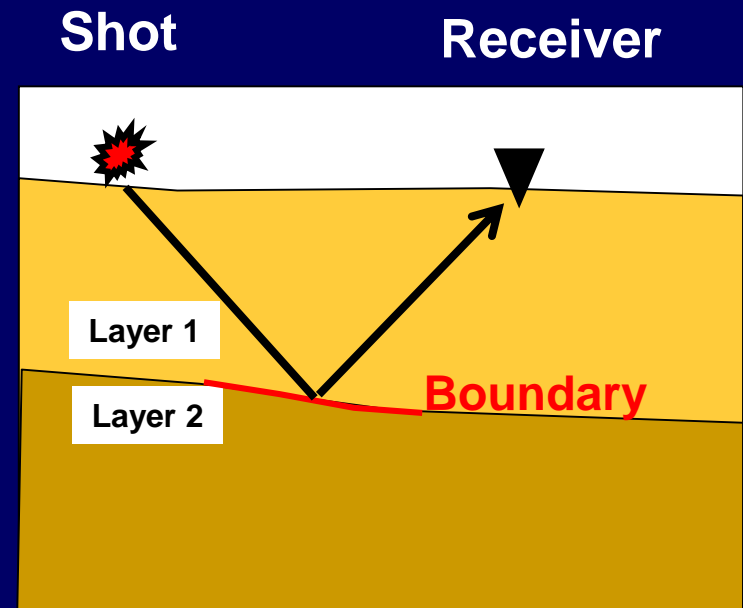
- Positive # on the tape
- Displayed as a Peak

SEG = Society of Exploration Geophysics

What Causes Reflections?

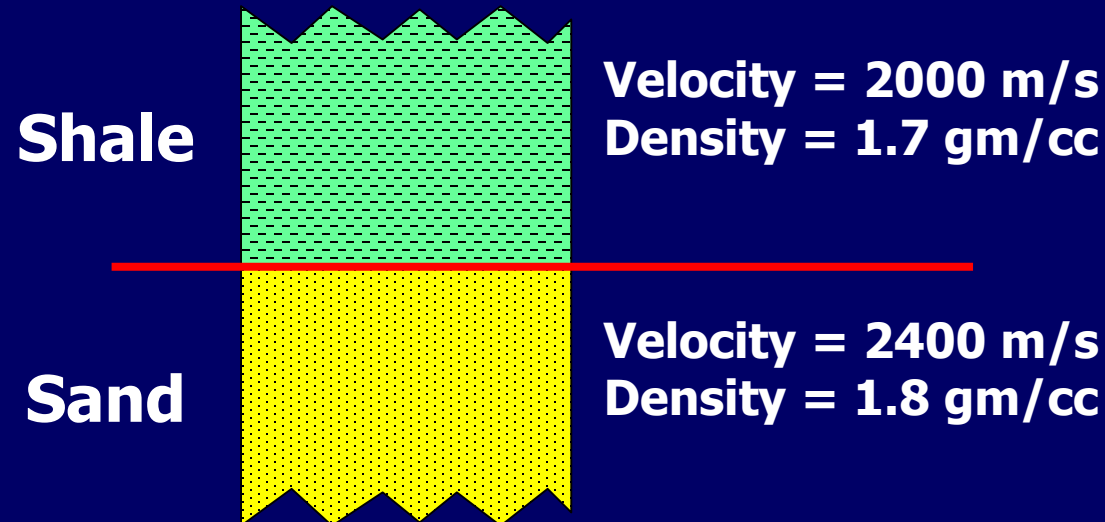


- Any interface between bodies with different acoustic properties
- Acoustic properties define **Impedance (I)**, in which
 $I = \text{velocity} * \text{density}$



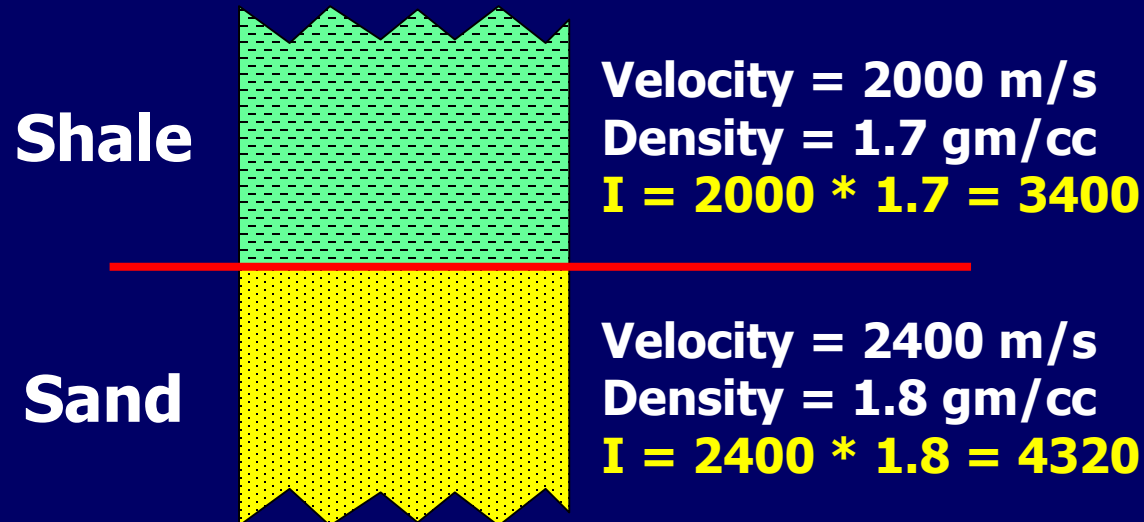
- Small change in impedance – small reflection
- Large change in impedance – large reflection

Seismic Interface



$$\text{Reflection Coefficient} = \frac{I_{\text{below}} - I_{\text{above}}}{I_{\text{below}} + I_{\text{above}}} = \underline{\hspace{2cm}} =$$

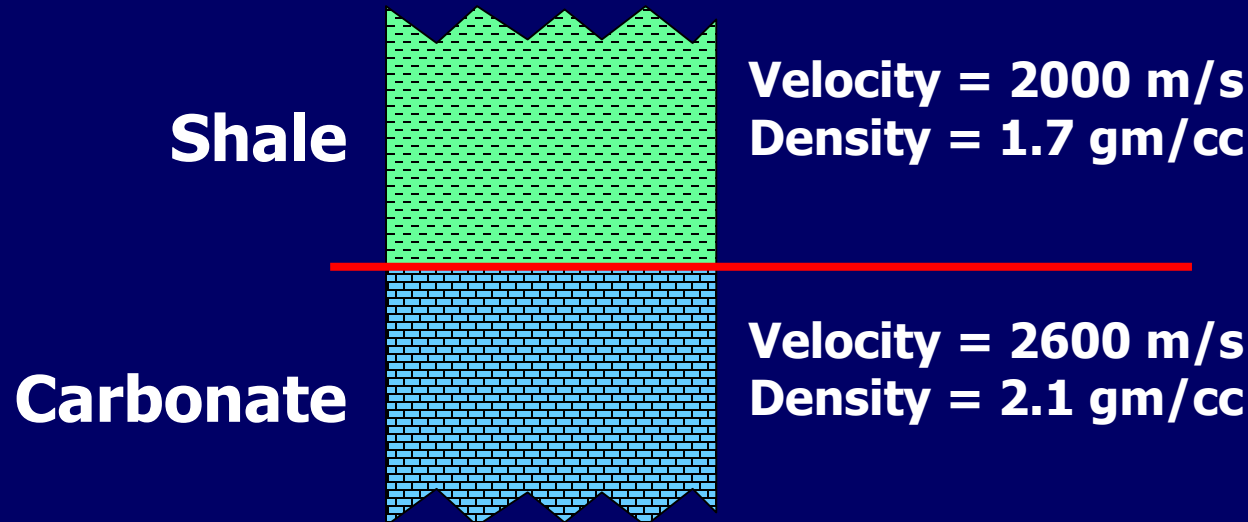
Seismic Interface



$$\text{Reflection Coefficient} = \frac{I_{\text{below}} - I_{\text{above}}}{I_{\text{below}} + I_{\text{above}}} = \frac{4320 - 3400}{4320 + 3400} = 0.119$$

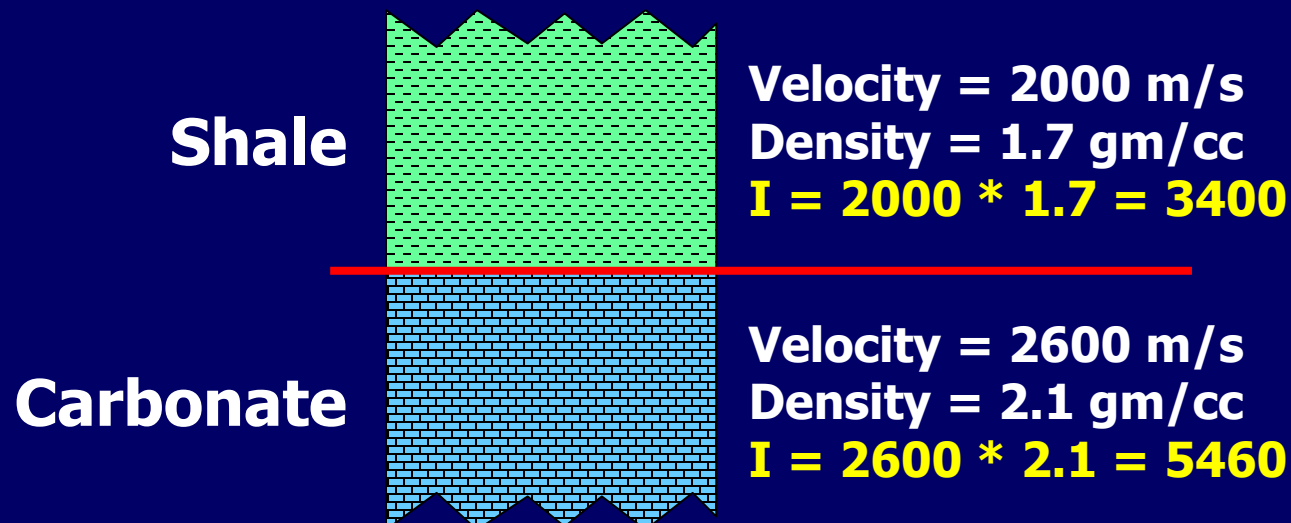
Of the incident energy, 12% is reflected, 88% is transmitted

Seismic Interface



$$\text{Reflection Coefficient} = \frac{I_{\text{below}} - I_{\text{above}}}{I_{\text{below}} + I_{\text{above}}} = \underline{\hspace{2cm}} =$$

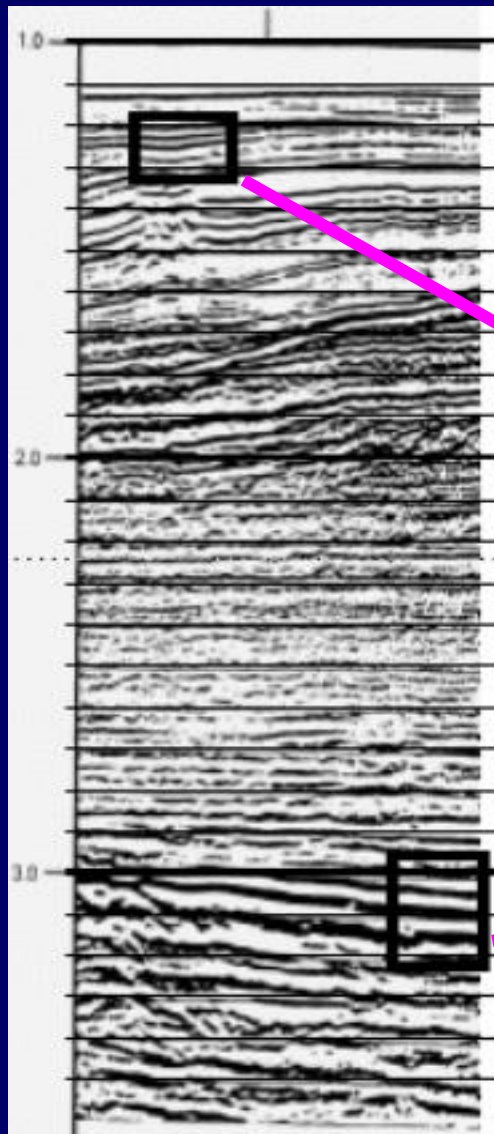
Seismic Interface



$$\text{Reflection Coefficient} = \frac{I_{\text{below}} - I_{\text{above}}}{I_{\text{below}} + I_{\text{above}}} = \frac{5460 - 3400}{5460 + 3400} = 0.232$$

Of the incident energy, 23% is reflected, 77% is transmitted

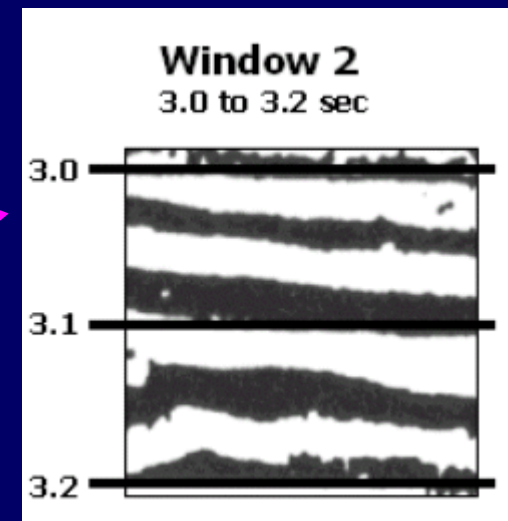
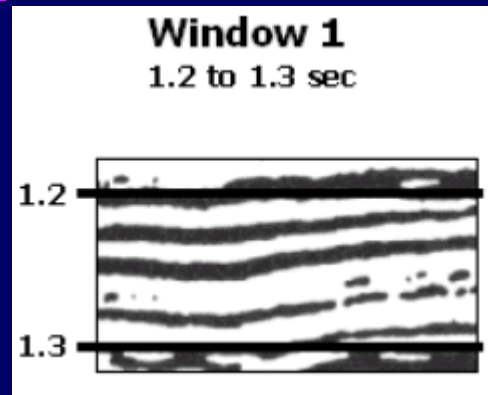
Exercise 6b: Frequency & Wavelength



$$F_{\text{apparent}} = \# \text{ cycle} / \text{time interval}$$

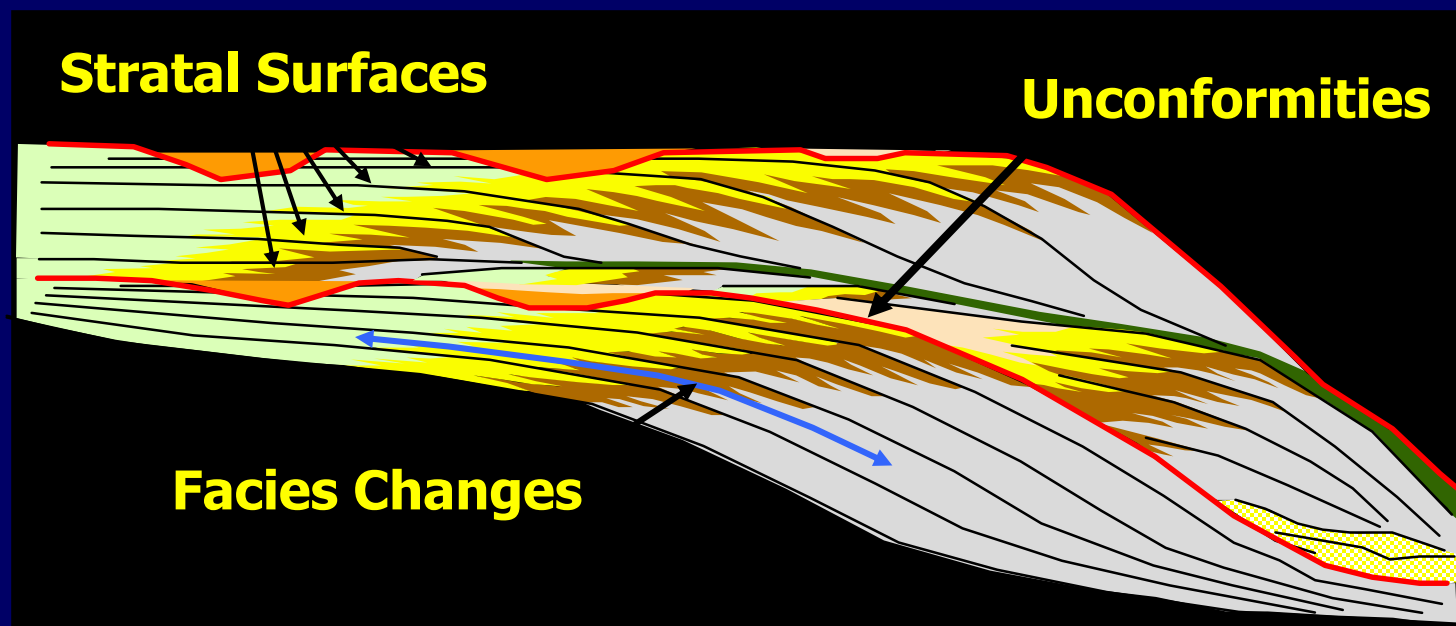
$$F_{\text{dominant}} = (2 * F_{\text{apparent}}) / \pi$$

$$\lambda = \text{Velocity} / F_{\text{dominant}}$$



Seismic Data & Stratal Surfaces

- Seismic reflections parallel stratal surfaces
- Reflection terminations mark unconformities
- Changes in reflection character indicate facies changes



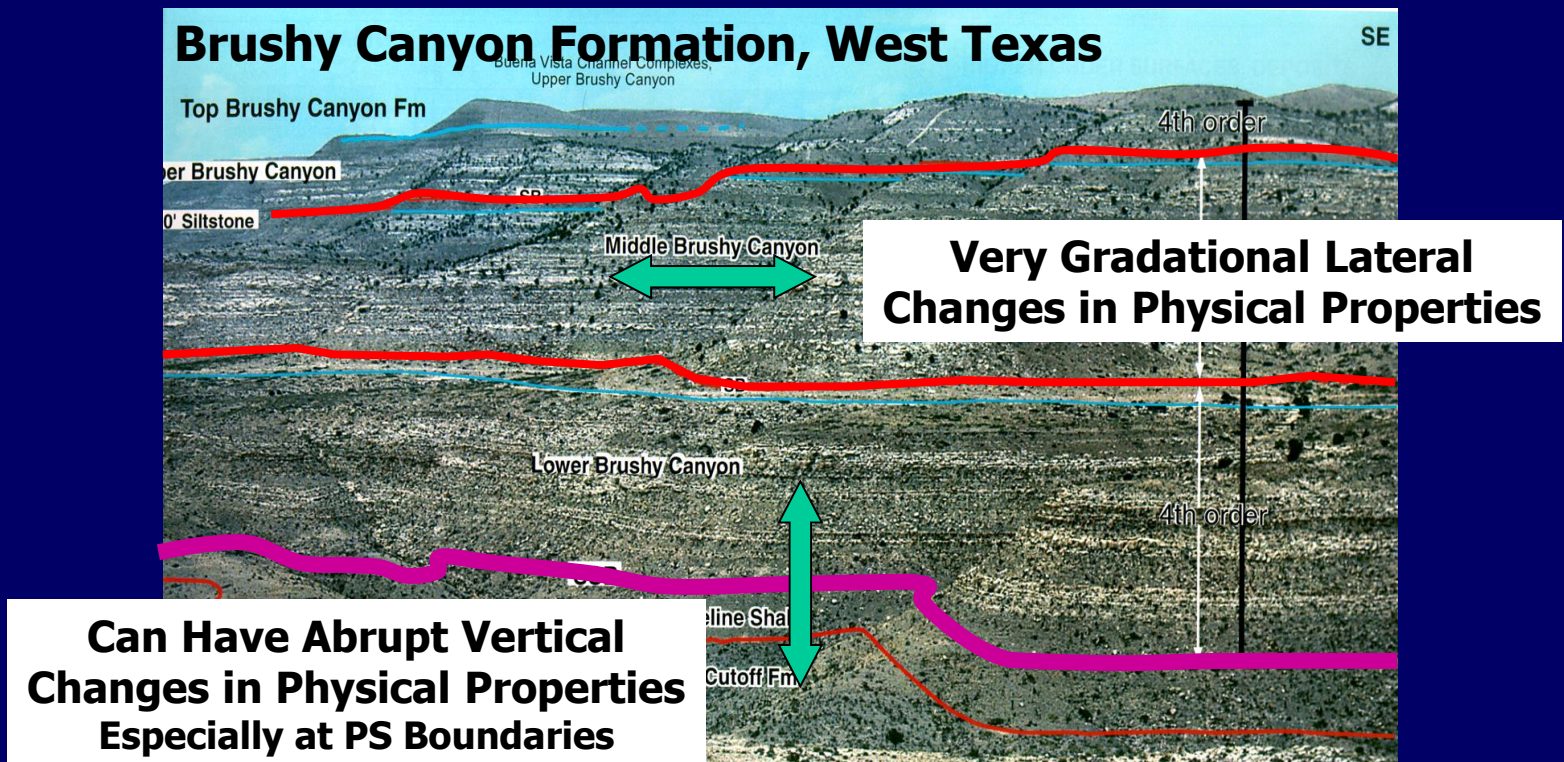
Fluvial Incised Valley Fill	Coastal Plain	Foreshore/Upper Shoreface
Lower Shoreface - Offshore	Slope - Basin	Submarine Fan
Estuarine	Condensed Interval	

Why Stratal Surfaces?

Recall: Reflections are generated where there is a change in acoustic properties ($I = \rho v$)

Consider: Where can there be sharp changes in impedance?

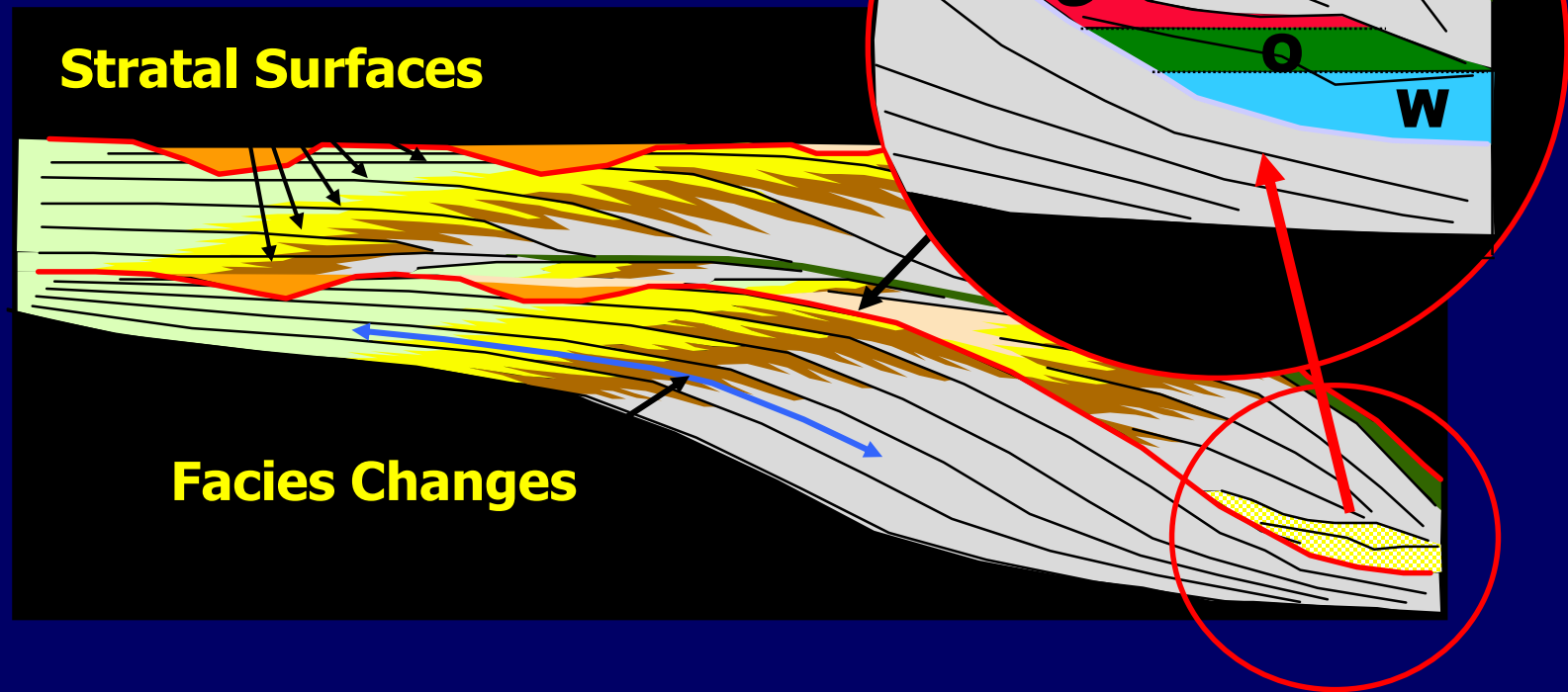
- horizontally as lithofacies change?
- vertically across stratal boundaries?



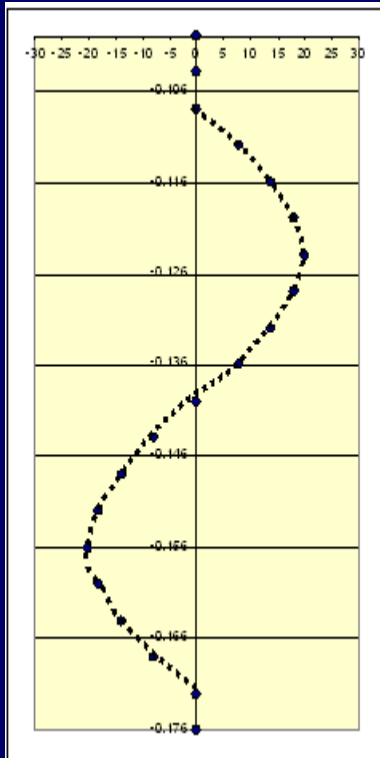
Not Every Reflection is Strata!

There are other seismic reflections out there that may not be stratigraphic in origin

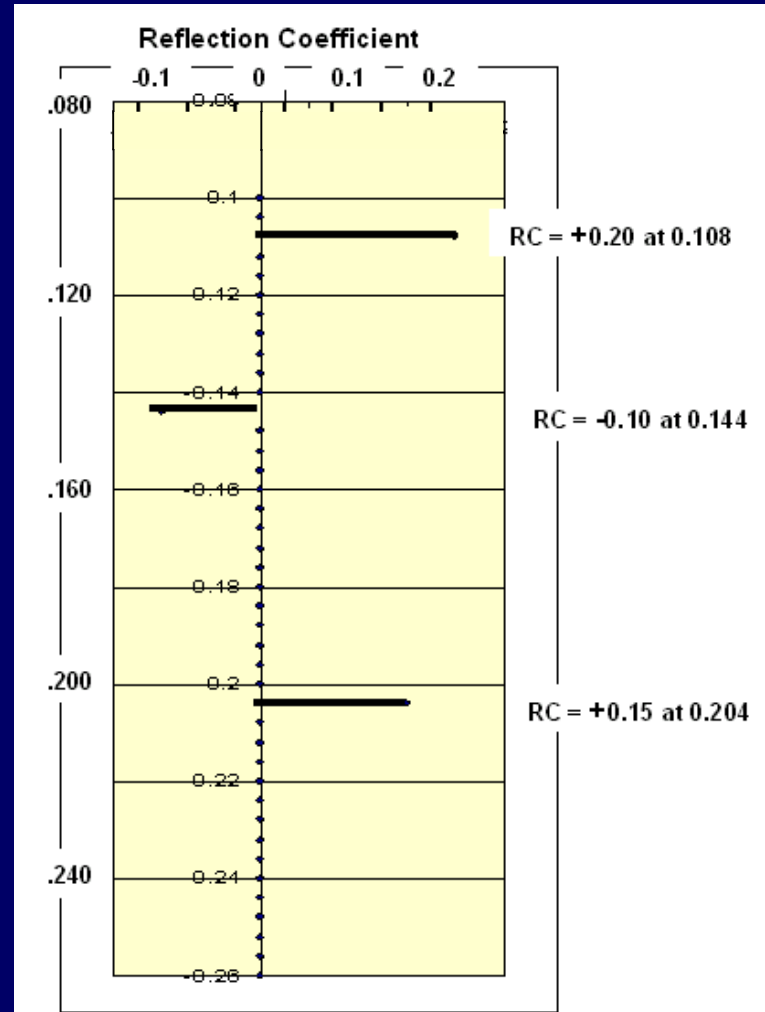
- Fluid Contacts
- Fault Planes
- Multiples
- Others



Exercise 6c: A Synthetic Trace



The Pulse



3 Ref. Coeff.